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Panel Discussion of Gravity Waves

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Saturation Theories

1. "Linear instability" theory

- Dewan + Good ('86), Smith et al. ('87)

"Saturated cascade"

- Dewan ('91)

positives:

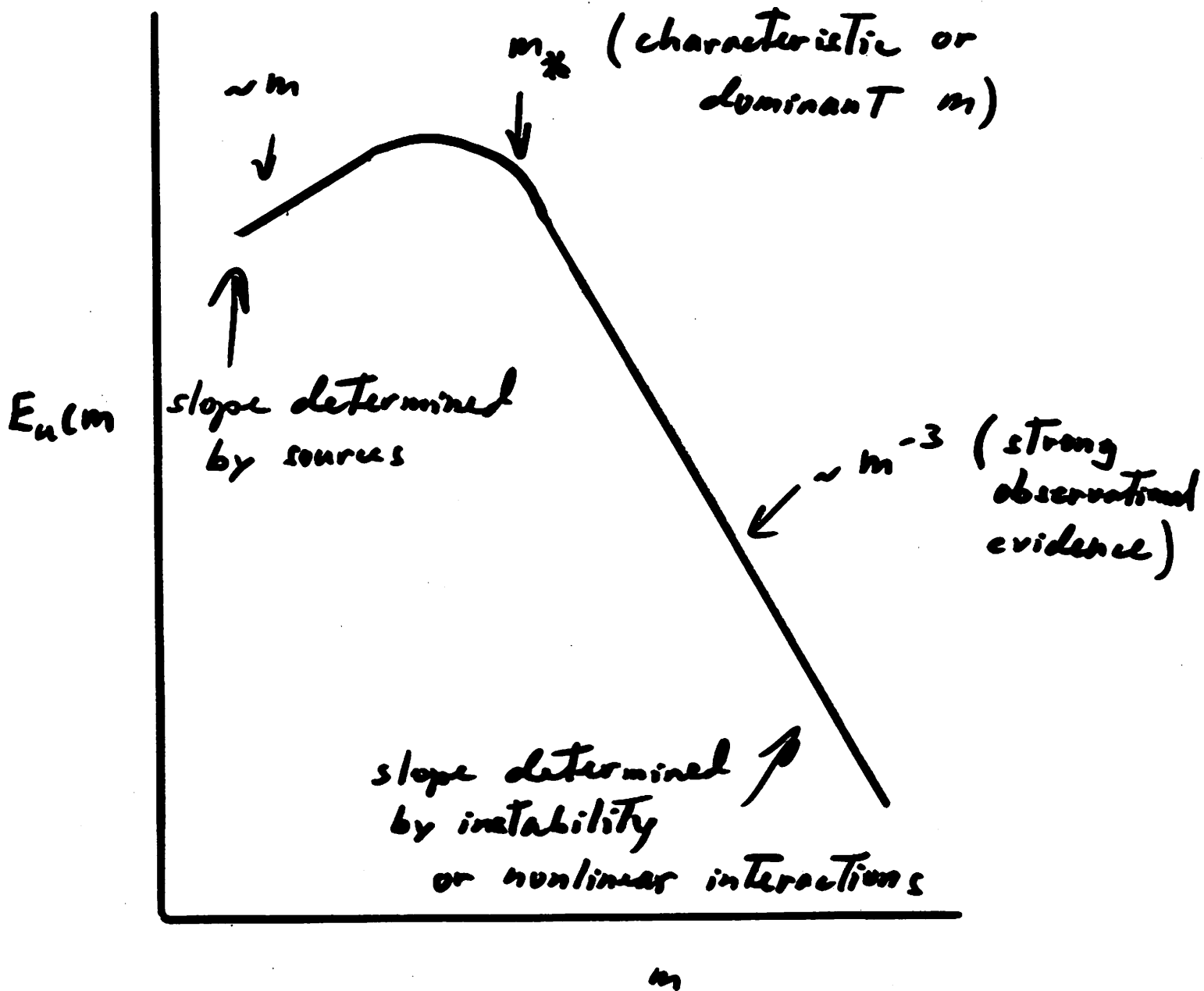
- clear observ. evidence for instab. constraints on wave amps. (Hodges, '67)
- accounts for N^2 variations, scale changes, decreasing m_x , $E(z)$ vars. naturally

negatives:

- ampl. constraints ad hoc
- packet bandwidth arbitrary
- does not include wave-wave int. effects

specific failing:

- stratospheric scaling of ampl. with N^2 does not account for wave filtering by shear and N^2 reductions



2. Doppler spread Theory

- Hines ('91, ...)

positives:

- conceptually simple + appealing
- has led to GW parameterization with some nice attributes

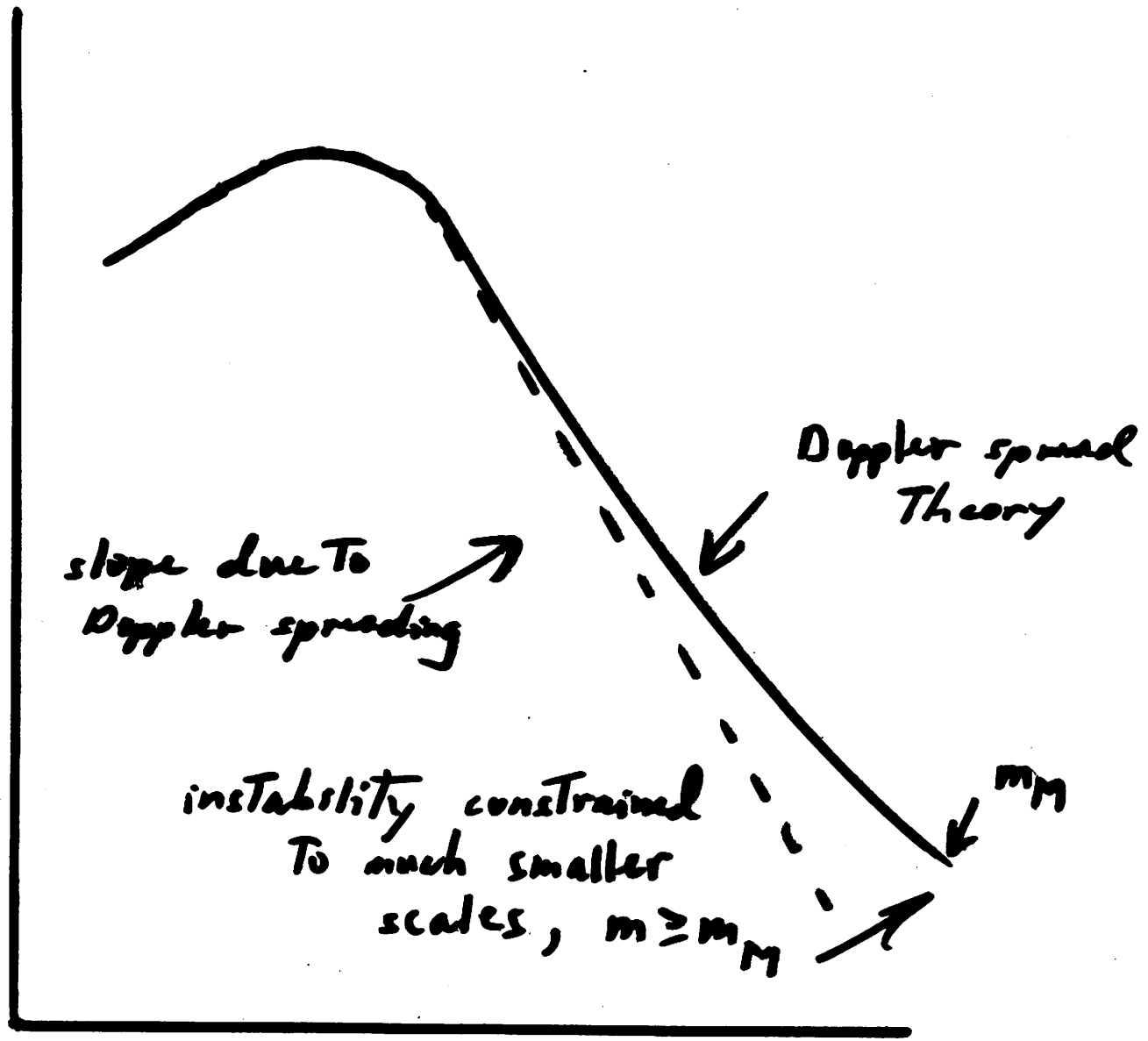
negatives:

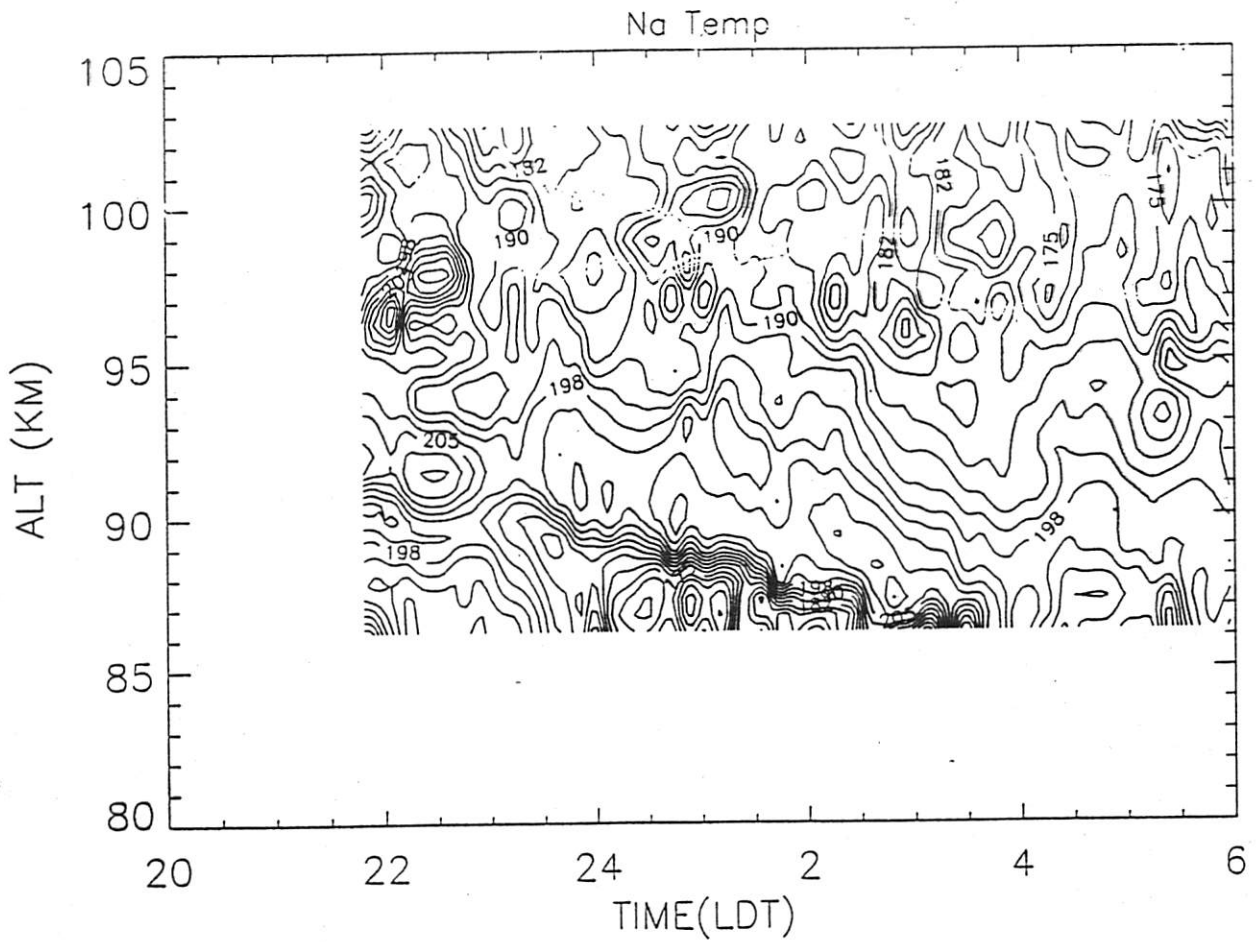
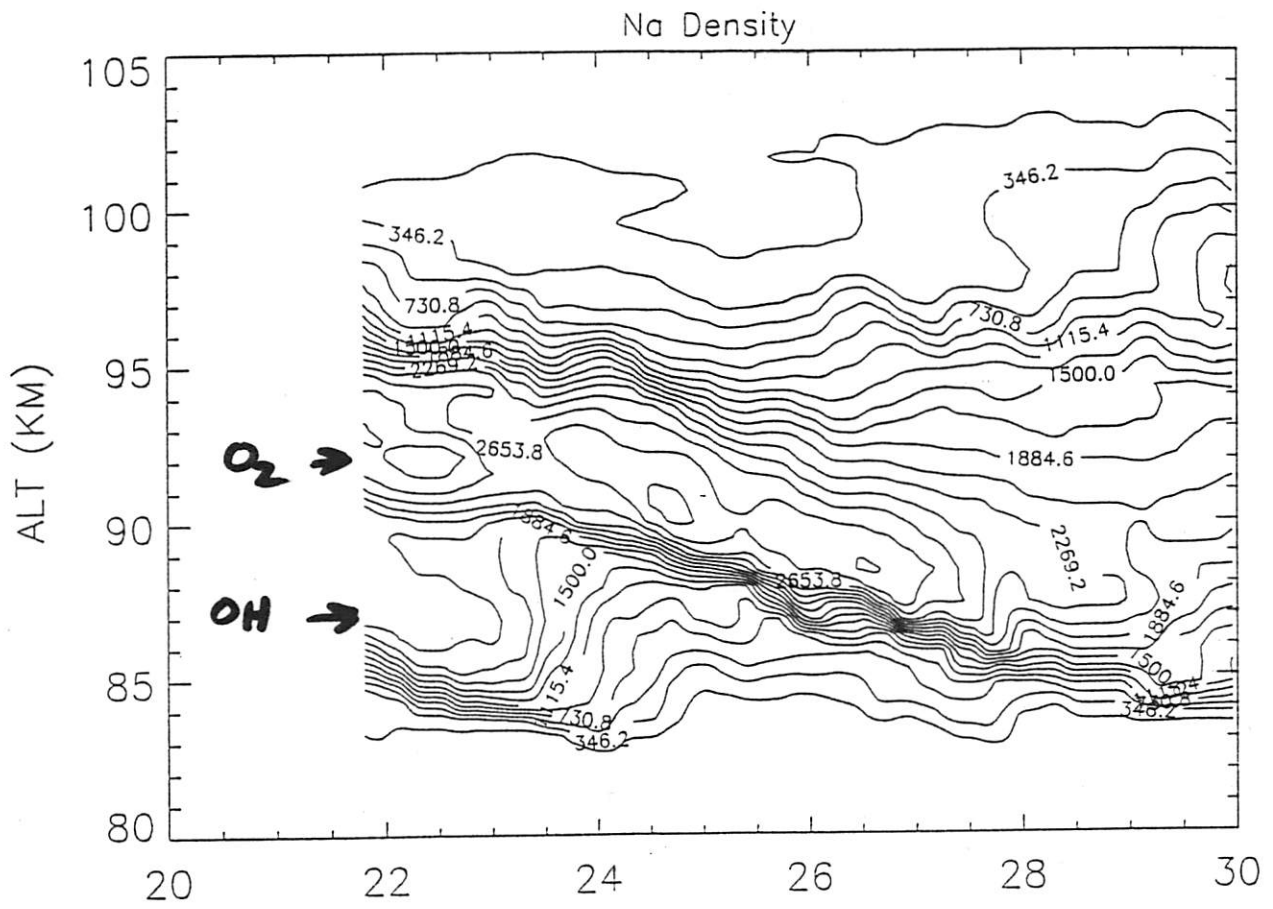
- fails to account for observed "Tail" spectrum, $\sim m^{-2.7}$ or shallower
- fails to account for wave dissipation at observed scales ($m \ll m_p$)

major failing:

- primary approximation of OS Theory is not valid
- fails consistency check with ray theory (Eckermann, '97, JAS, in press)

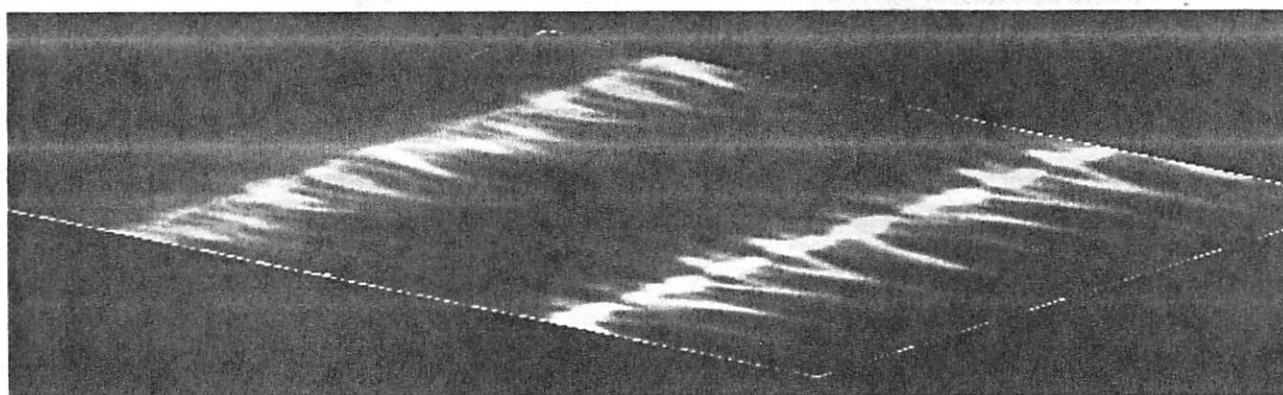
Doppler spread spectrum





CORN Hecht et al. ('97)

Geophysical Research Letters



OCTOBER 8, 1993

Volume 20 Number 19

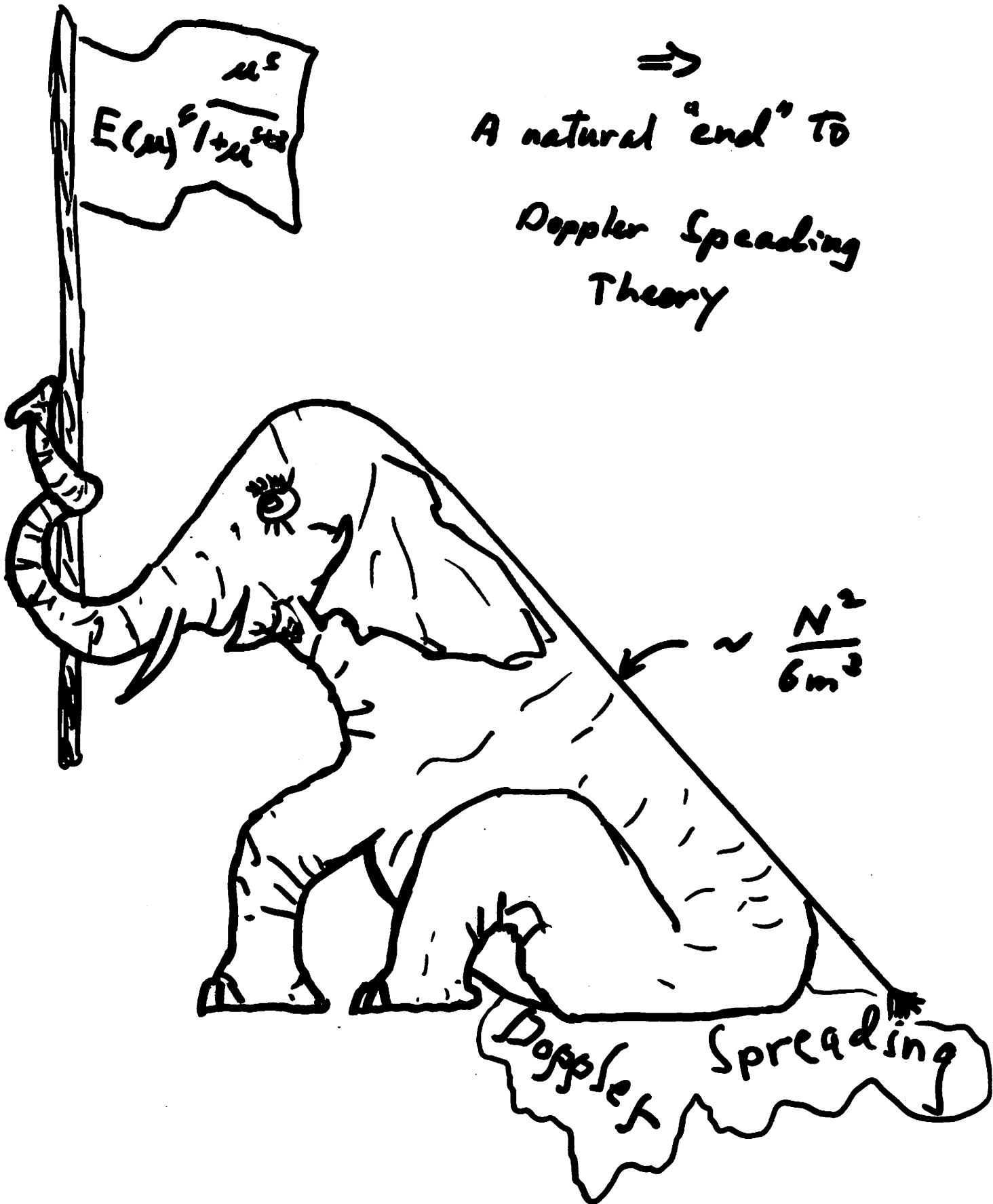
AMERICAN GEOPHYSICAL UNION

The "weight" of evidence



A natural "end" to

Doppler Spreading
Theory



3. Nonlinear diffusion + variants

- Weinstock ('90), Zhu ('94), Medvedev + Klaassen ('95)
- Gardner ('94) - "diffusive filtering"

positives:

- account for partial advection NL's
- predicts correct spectral shapes w/in amplitude factor for $E_{\omega}(m)$

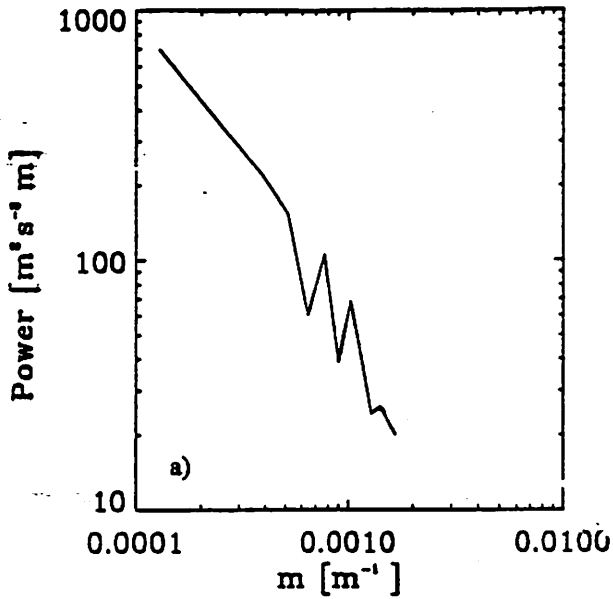
negatives:

- "NL" theories neglect certain NL ints.
- "diffusive filtering" neglects wave-wave ints
- damping mechanism ad hoc, speculative

major failing:

- Testable + incorrect predictions for $E_{\omega}(m) \sim m^p$ for $p=2$ (ω^{-p}) by diffusive filtering theory

Average m-Spectrum
 80-87.2 km; 8:42 - 11:29 UT
 Integral = $0.37 \text{ m}^2 \text{ s}^{-2}$



Average m-Spectrum
 83-89.3 km; 12:43 - 13:42 UT
 Integral = $0.92 \text{ m}^2 \text{ s}^{-2}$

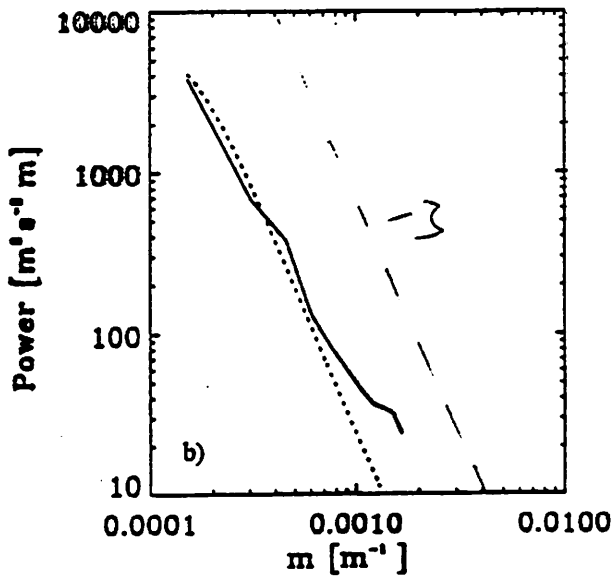


Figure 5. Vertical wavenumber spectra for altitudes from 80 to 87.2 km averaged (a) from 0842 to 1129 UT and from (b) 83 to 89.3 km averaged from 1243 - 1342 UT. The spectra are representative of ducted and propagating wave environments, respectively. Note that the integrated variances are much less than those obtained from the time series in each case because of the mean and linear trend removal. The canonical vertical wavenumber spectrum having the same variance is shown with a dashed line for reference in Figure 5b.

Fritts & Hoppe ('95, JGR)

SOUSY VHF - vertical velocity

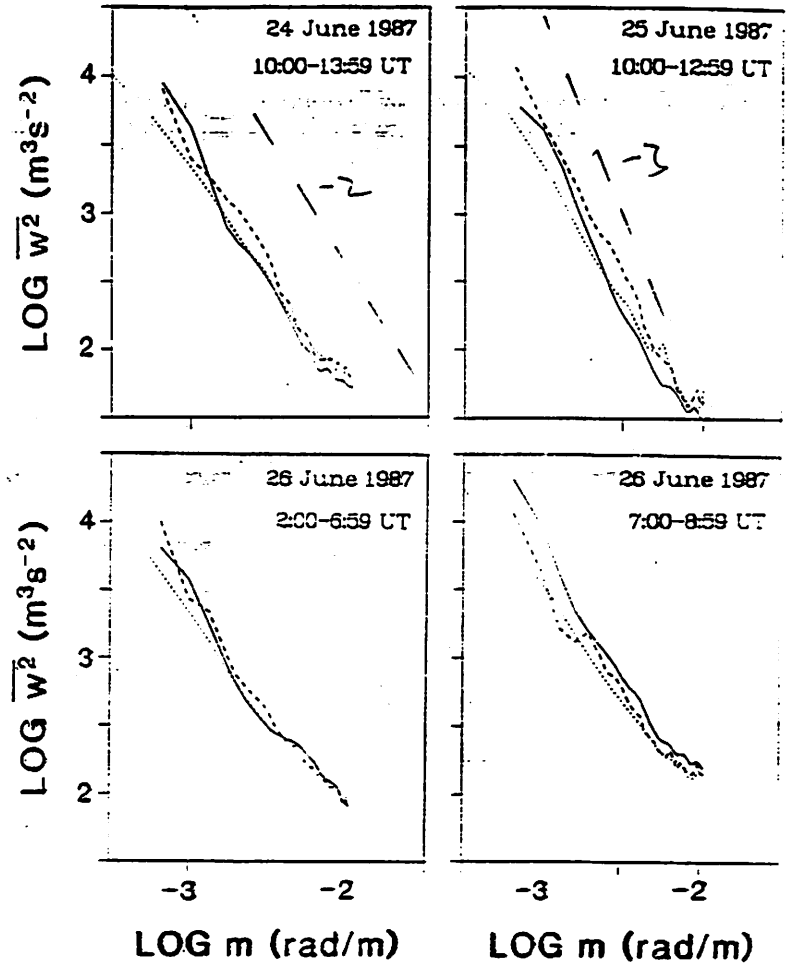


Fig. 11. SOUSY vertical wavenumber spectra averaged for four periods of 2- to 5-h duration using an FFT with a cos-taper (—), an FFT with a box window and linear ramps (---) and MEM with no tapering (···). Slopes at lower wavenumbers approach -3 . At higher wavenumbers, the spectra are likely influenced by measurement errors and proximity to L_B .

Fritts et al. ('90, JATP)

Summary:

- All present Theories are
 - ill-posed, incomplete, or incorrect in significant ways
 - Yet all Theories also contain aspects which have helped to illuminate the physical processes we must describe more completely
 - Theoretical difficulties \Rightarrow
 - The full NL problem including
 - spectral (wave-wave) interactions
 - local wavefield instabilitiesis analytically intractable
- \Rightarrow we will likely need to rely on numerical methods to address the interplay between these processes more completely.